

1 **Wind induced algal migration manipulates sediment denitrification N-loss patterns in**  
2 **shallow Taihu Lake, China**

3 Wenqing Shi<sup>a,b</sup>, Lin Zhu<sup>a</sup>, Bryce Van Dam<sup>c</sup>, Ashley R. Smyth<sup>d</sup>, Jianming Deng<sup>e</sup>, Jianzhou<sup>e</sup>,  
4 Gang Pan<sup>f</sup>, Qitao Yi<sup>g</sup>, Jianghua Yu<sup>a</sup>, Boqiang Qin<sup>e,\*</sup>

5 <sup>a</sup>Jiangsu Collaborative Innovation Center of Atmospheric Environment and Equipment  
6 Technologies, Jiangsu Key Laboratory of Atmospheric Environmental Monitoring & Pollution  
7 Control, School of Environmental Science & Engineering, Nanjing University of Information  
8 Science & Technology, Nanjing 210044, China

9 <sup>b</sup>Hubei Provincial Key Laboratory of River Basin Water Resources and Eco-Environmental  
10 Sciences, Changjiang River Scientific Research Institute, Wuhan 430010, China

11 <sup>c</sup>Institute of Carbon Cycles, Department of Fluxes across Interfaces, Helmholtz-Zentrum  
12 Hereon, Geesthacht, 21502, Germany

13 <sup>d</sup>Soil and Water Sciences Department, Tropical Research and Education Center, University of  
14 Florida, Homestead, FL 33031, USA

15 <sup>e</sup>Taihu Lake Laboratory Ecosystem Station, State Key Laboratory of Lake Science and  
16 Environment, Nanjing Institute of Geography and Limnology, Chinese Academy of Sciences,  
17 Nanjing 210008, China

18 <sup>f</sup>School of Animal, Rural and Environmental Sciences, Nottingham Trent University,  
19 Brackenhurst Campus, Southwell NG25 0QF, United Kingdom

20 <sup>g</sup>School of Civil Engineering, Yantai University, Yantai 264005, China

21 \*Corresponding author: Boqiang Qin ([qinbq@niglas.ac.cn](mailto:qinbq@niglas.ac.cn))

22 **Abstract**

23 Driven by winds, the distribution of algae is often noticeably patchy at kilometer scales in  
24 shallow lakes. The decomposition of the settled algal biomass may affect N biogeochemical  
25 cycles and thereby nitrogen (N) loss in sediments. In this study, we investigated sediment  
26 denitrification N-loss patterns along algal migration pathway in Taihu Lake, a shallow and  
27 eutrophic lake in China, and found that wind-induced algal migration in the overlying water  
28 manipulated the temporal and spatial patterns of denitrification N-loss in sediments. A N loss  
29 hotspot in sediments was created in the algae concentrated zone, where N loss was, however,  
30 temporarily inhibited during algal bloom seasons and generally exhibited a negative  
31 relationship with algal biomass. In the zone where algae have left, sediment N loss rate was  
32 relatively low and positively correlated with algal biomass. The decay of algal biomass  
33 generated organic carbon and created anoxia, favoring denitrification, while excessive algal  
34 biomass could deplete oxygen and inhibit nitrification, causing nitrate limitation for  
35 denitrification. Piecewise linear regression analysis indicated that algal biomass of *Chl-a* >  
36 73.0  $\mu\text{g/L}$  in the overlying water could inhibit denitrification N-loss in sediments. This study  
37 adds to our understanding of N biogeochemical cycles in shallow eutrophic lakes.

38 **Keywords:** denitrification, algal migration, nitrogen loss, sediment, shallow lake

## 39 **1. Introduction**

40 Human activities have greatly increased nitrogen (N) loading into the environment (Finlay et  
41 al., 2013; Horst et al., 2014; Paerl et al., 2011), with damaging effects including eutrophication  
42 and subsequent odoriferous algal blooms, water quality degradation and biodiversity reduction  
43 (Bianchi et al., 2010; Lewis Jr et al., 2011; Porter et al., 2013). In lakes, there are many nitrogen  
44 removal pathways, among which microbial nitrification (NF) and denitrification (DNF) in  
45 sediments play a key role in alleviating the negative impacts of excessive N by converting  
46 reactive N to gaseous forms and permanent removal them from lakes (Chen et al., 2012; Xia et  
47 al., 2016; Xia et al., 2009). In this process, NF is responsible for the oxidation of ammonium  
48 ( $\text{NH}_4^+\text{-N}$ ) to nitrate ( $\text{NO}_3^-\text{-N}$ ), which is in turn a substrate for DNF (Gao et al., 2012; Kessler  
49 et al., 2013; Xia et al., 2016), and any perturbation on the coupled NF-DNF may alter N loss  
50 in the sediment (Gao et al., 2012; Kessler et al., 2013; Small et al., 2014). For instance, anoxia  
51 in permeable sediments inhibits NF and thereby breaks the coupled NF-DNF, leading to limited  
52 N loss in the hyporheic zone (Kessler et al., 2013; Kessler et al., 2012).

53 In algae dominated shallow lakes, the DNF-NF in sediments is mediated by the settling of algal  
54 biomass from the overlying water. Dissolved inorganic N in the overlying water is assimilated,  
55 converted to organic N, and sunk to bed sediments by algae. The degradation of organic N  
56 releases  $\text{NH}_4^+\text{-N}$ , which is subsequently oxidized to  $\text{NO}_3^-\text{-N}$  through NF and eventually  
57 converted back to  $\text{N}_2$  through DNF (Kuypers et al., 2018). Meanwhile, this deposition of algae-  
58 derived organic carbon results in oxygen consumption, and a transition towards anoxia, a  
59 requirement for heterotrophic denitrifying bacteria, and therefore DNF (Conley, 2012; Wang

60 et al., 2016). Thus, sediment N loss, through DNF, is correlated with the accumulation of algal  
61 biomass in sediments (Chen et al., 2016). While anoxia/hypoxia favors DNF, it can also inhibit  
62 aerobic NF, causing nitrate limitation to DNF. Zhu et al. (2020) found that, when algal biomass  
63 was experimentally increased by a factor of 5 and 10, total sediment DNF rates were decreased  
64 by 32% and 79%, respectively. Such dense accumulations of algal biomass regularly occur in  
65 the shallow lakes, and are often caused by wind-driven algal redistribution (Marcé et al., 2010).  
66 This is because many bloom-forming algal species (e.g., *Microcystis* sp.) contain intracellular  
67 gas vesicles and are positively buoyant, making them prone to concentrate at the downwind  
68 shore (Cyr, 2017; Deng et al., 2016; George and Edwards, 1976). Cyr (2017) demonstrated that  
69 algal biomass was higher at the downwind sites than those at the upwind sites on windy days,  
70 and the magnitude of this difference increased linearly with increasing wind speed. Dense algal  
71 mats are also observed in the downwind northwestern bays following prevailing southeast  
72 winds in Taihu Lake, a eutrophic lake in China (Deng et al., 2016; Qin et al., 2010; Zhang et  
73 al., 2015). We hypothesize that, in shallow eutrophic lakes, (i) wind-driven algal migration in  
74 the overlying water manipulates DNF N-loss patterns in sediments; (ii) sediment DNF N-loss  
75 is relatively weak in the zone where algae have left, while it is strong but will be temporally  
76 inhibited during algal bloom seasons in the algae concentrated zone; (iii) there exists an algal  
77 biomass threshold, and sediment DNF N-loss rate increases beyond the threshold, after which  
78 the rate decreases as algal biomass further increases.

79 To test the hypothesis, we investigated DNF ( $N_2$  gas production) along the algal migration  
80 pathway in Taihu Lake, and identify the impacts of algal biomass in the overlying water on

81 sediment DNF N-loss. These rate measurements were combined with a genetic investigation  
82 targeted at the microbial pathways responsible for sediment NF and DNF. The objective of this  
83 study was to explore the impacts of wind-driven algal migration on sediment N biogeochemical  
84 cycling in a shallow and eutrophic lake.

## 85 **2. Methods and materials**

### 86 **2.1. Study area**

87 Taihu Lake, the 3rd largest freshwater lake (~2,400 km<sup>2</sup>) in China. Taihu Lake has a mean  
88 depth of 1.9 m, a volume of 4.4 billion m<sup>3</sup> and a mean water retention time of 284 days (Qin et  
89 al., 2019). Taihu Lake is located in the Yangtze River delta (the lake center coordinates: 31°10'0"  
90 N, 120° 9'0" E), which is the most rapidly developing region of China (Figure 1A). The Taihu  
91 Basin accounts for only 0.4% of China's land area, but contributes to 11% of its Gross  
92 Domestic Product (GDP) (Xu et al., 2010). Taihu Lake is a key drinking water, fishing and  
93 tourism resource for the region, but also serves as a depositional basin for urban, agricultural  
94 and industrial segments of the local economy. With recent increases in pollutant (N & P)  
95 loadings, Taihu Lake has experienced accelerated eutrophication and severe algal blooms in  
96 recent decades. These blooms have caused serious environmental, economic and societal  
97 problems, including a threat to drinking water supplies for ~10 million residents. In May 2007,  
98 a massive algal bloom overwhelmed the lake's water treatment plants, leaving more than two  
99 million people without drinking water for a week (Paerl et al., 2011; Qin et al., 2019).

### 100 **2.2. Field surveys**

101 Field samplings were conducted along two hypothetical algal migration pathways from the lake  
102 center to two northwestern bays, Meiliang Bay and Gonghu Bay (Figure 1B), through which  
103 different amounts of algal biomass are transported. This will be beneficial to quantifying the  
104 relationship between algal biomass and DNF rate. There were 6 sampling sites in each corridor,  
105 from lake center to each bay, with a shared site at the lake center (11 sites in total), among  
106 which the shared site was used to represent lake center, and the site at the end of each corridor  
107 was used to represent lake bay, respectively. The distance between two neighboring sites was  
108 about 6 km. Six field surveys were conducted in Jan, Apr, Jun, Jul, Aug and Nov (algal bloom  
109 period: June-August) of 2018. On the sampling day, surface water (0–20 cm) and sediment  
110 were sampled in triplicate at each site using a stainless-steel bucket and an Ekman grab sampler,  
111 respectively. A water sample of 500 mL was preserved on ice for the analyses of chlorophyll a  
112 (*Chl-a*) and various N species, including total N (TN),  $\text{NH}_4^+$ -N,  $\text{NO}_3^-$ -N and nitrite ( $\text{NO}_2^-$ -N);  
113 after homogenized completely, a sediment sample of 50 g was frozen in liquid N for microbial  
114 analysis. Intact sediment cores (9 cm in diameter and 50 cm in height) were also carefully  
115 collected at each site in triplicate using a Jenkin corer for measuring sediment N loss rate as  
116 described in the following section. N species were analyzed using the Monitoring Analysis  
117 Method of Water and Waste Water (Ministry of Environmental Protection of China, 2002).  
118 Before the analyses of dissolved N species, water samples were filtered through Whatman  
119 GF/F membrane filters. *Chl-a* was determined using the hot ethanol method according to  
120 Pápista et al. (2002).

### 121 **2.3. Sediment net $\text{N}_2$ flux measurements**

122 Sediment net N<sub>2</sub> flux was measured using intact sediment cores (Figure S1) (Heiss et al., 2012).  
123 Immediately after collection, sediment cores were transported to the lab at Taihu Lake  
124 Laboratory Ecosystem Research Station. When algal accumulation occurs, light hardly reaches  
125 bed sediments under dense algal canopies (Thomsen et al., 2012), these sediment cores were  
126 therefore incubated in the dark in a water bath at the field temperature. Sediment cores were  
127 carefully sealed with a gas-tight rubber lid without air bubbles, and left to stand for 10 min to  
128 ensure equilibrium according to our preliminary experiments. Considering oxygen and  
129 substrates may change in the sealed system, potentially affecting the measured value, we  
130 collected samples in a short time to maintain DO above 2 mg/L. Water samples were collected  
131 every 5 min over a 20-min period and stored into a 12-mL pre-evacuated Exetainer® vial (839  
132 W, Labco, UK) after adding 0.2 µL saturated HgCl<sub>2</sub> solution. Net DNF rates were calculated  
133 based on the linear rate of N<sub>2</sub> accumulation in the overlying water as a function of time. The  
134 concentration of dissolved N<sub>2</sub> was measured by the N<sub>2</sub>/Ar method using a membrane inlet mass  
135 spectrometry system (Bay Instruments, Easton, MD, USA). These net N<sub>2</sub> flux measurements  
136 may represent the balance between N<sub>2</sub> fixation and DNF. However, sediment N fixation in  
137 these bays was estimated to be only 1.8% of sediment N loss via denitrification (Yao et al.,  
138 2018). We argue that these net N<sub>2</sub> flux measurements to be representative of potential DNF  
139 rates.

#### 140 **2.4. Microbial abundance analysis**

141 The abundances of nitrifiers and denitrifiers in sediments were analysed using quantitative  
142 polymerase chain reaction method (qPCR). There are many functional genes involved in NF

143 and DNF processes (Kuypers et al., 2018), of which *amoA*, *nirS*, *nirK* and *nosZ* were chosen  
144 as gene markers in this study. The *amoA* is a gene marker of ammonia-oxidizing bacteria,  
145 which are sensitive indicators of NF and play a more important role in the NF in Taihu Lake  
146 (Hou et al., 2015); *nirS*, *nirK* and *nosZ* are commonly used gene markers of denitrifiers, which  
147 encode nitric oxide reductase, nitrite reductase and nitrous oxide reductase, respectively (Hou  
148 et al., 2015; Li et al., 2019; Morales et al., 2010). Prior to analyses, sediment samples were  
149 stored at -80°C. DNA was extracted for qPCR amplification using a FastDNA Power-Max Soil  
150 DNA Isolation Kit (MP Biomedical, USA) according to the manufacturer's instructions. The  
151 detailed information for qPCR was presented in Text S1 in the supporting information.

## 152 **2.5. Statistical analysis**

153 Piecewise linear regression model was employed to investigate the impacts of algal biomass in  
154 the overlying water on sediment N loss. Piecewise linear regression can identify where the  
155 slope of a linear function changes and allow multiple linear models to be fitted to each distinct  
156 section (Turner et al., 2018; Zeng et al., 2019). To avoid the disturbances caused by temperature,  
157 data at the 11 sampling sites in July and August were used here, when water temperature  
158 showed no significant differences (Table S1). Piecewise linear regression analysis was  
159 conducted using MATLAB (MathWorks, USA). T-test was applied to test the differences  
160 between lake bay and the center using SPSS 22.0 (SPSS Inc., North Chicago, IL, USA). The  
161 level of statistical significance was set at  $P < 0.05$ .

## 162 **3. Results**

### 163 **3.1. *Chl-a* and TN in the overlying water**



164 In general, there was a remarkable increasing trend of algal biomass in the overlying water  
165 from the lake center to lake bays, especially during the algal accumulation season (Figure 2A).  
166 *Chl-a* increased from 8.7–15.6  $\mu\text{g/L}$  in lake center to 22.7–56.3  $\mu\text{g/L}$  in Meiliang Bay (M5)  
167 and 18.5–25.7  $\mu\text{g/L}$  in Gonghu Bay (G5) during the non-algal bloom period (Nov-Apr), and  
168 increased from 13.5–18.1  $\mu\text{g/L}$  in the lake center to 136.5–227.3  $\mu\text{g/L}$  in Meiliang Bay (M5)  
169 and 54.9–112.2  $\mu\text{g/L}$  in Gonghu Bay (G5) during the algal bloom period (Jun-Aug). Similarly,  
170 TN concentration in the bays was higher than in the lake center, especially during the algal  
171 bloom period. TN in Meiliang Bay (M5) and Gonghu Bay (S5) was 1.75–1.96 and 1.77–2.33  
172 mg/L, respectively, during the non-algal bloom period. TN in these bay sites then reached  
173 maximal values of 3.83 and 2.34 mg/L during the algal bloom period, while TN in the lake  
174 center remained relatively consistent at 1.05–2.12 mg/L throughout the year (Figure 2B).

### 175 **3.2. Sediment N loss rate**

176 Net DNF rates in the lake bays was remarkably higher than at the lake center, and also showed  
177 different temporal patterns than trends at the lake center (Figure 3). In the lake bays, sediment  
178 N loss rate was higher during the non-algal bloom period than the algal bloom period, while,  
179 in the lake center, the rate was slightly lower during non-algal bloom period than the algal  
180 bloom period. Sediment N loss rate reached 2.22–2.51 and 1.11–1.26  $\text{mg/m}^2\text{h}$  during the non-  
181 algal bloom period and decreased to 0.34–0.67 and 0.66–1.06  $\text{mg/m}^2\text{h}$  during the algal bloom  
182 period in the Meiliang Bay and Gonghu Bay respectively. In contrast, sediment N loss rate in  
183 the lake center was 0.09  $\text{mg/m}^2\text{h}$  during the non-algal bloom period and increased to a  
184 maximum of 0.42  $\text{mg/m}^2\text{h}$  during the algal bloom period (Figure 3).

### 185 **3.3. Relationship between N loss rate and algal biomass**

186 Sediment N loss rate as a function of time showed an opposite pattern to algal biomass in the  
187 overlying water in Meiliang Bay and Gonghu Bay (Figure 4B), but a similar pattern with algal  
188 biomass in the lake center (Figure 4A). Correlation analysis indicated that sediment N loss rate  
189 exhibited a negative relationship with algal biomass in the overlying water in Meiliang Bay ( $r^2$   
190 = 0.41) and Gonghu Bay ( $r^2 = 0.66$ ) (Figure 4D), but a positive relationship with algal biomass  
191 in the lake center ( $r^2 = 0.70$ ) (Figure 4C). Piecewise linear regression analysis was further  
192 conducted, indicating that sediment N loss rate increased beyond a threshold of 73.0  $\mu\text{g/L}$ , after  
193 which N loss rate decreased as algal biomass further increased (Figure 5).

### 194 **3.4. Relationship between functional microbes and algal biomass**

195 The abundances of nitrifiers and denitrifiers in sediments exhibited similar relationships with  
196 algal biomass in the overlying water (Figure 6). In both cases, the abundance remarkably  
197 increased at low levels of algal biomass and then gradually decreased as algal biomass further  
198 increased. The decreased microbe abundances mainly occur in algae concentrated lake bays.

## 199 **4. Discussion**

200 In lakes, wind can cause waves and subsequent advective movement of superficial water  
201 masses (Marce et al., 2007). When wind interacts with the lake surface, there is an energy  
202 transfer from the air towards the water, as mediated by wave-induced friction at the air-water  
203 interface. In shallow lakes, there is little space for wave energy to dissipate, and waves often  
204 have high turbulent kinetic-energy, initiating passive ‘migration’ of algal biomass (Wüest et  
205 al., 2000). The distribution of algae is noticeably patchy at kilometer scales in shallow

206 eutrophic lakes, in part due to this factor (Deng et al., 2016; Huang et al., 2015). In Taihu Lake,  
207 there is a subtropical monsoon climate with prevailing southeast winds, continuously  
208 transporting algal biomass by this mechanism towards downwind, northwestern bays (Deng et  
209 al., 2016). Accordingly, we detected a remarkable increasing trend of algal biomass in the  
210 overlying water from the lake center to Meiliang and Gonghu Bay during the algal  
211 accumulation season (Jun-Aug) (Figure 2A).

212 By algae, dissolved inorganic N ( $\text{NH}_4^+\text{-N}$  and  $\text{NO}_3^-\text{-N}$ ) in the overlying water is assimilated,  
213 converted to organic N, and eventually sunk to lake bed sediments. This transport of algal  
214 biomass results in the re-location of large amounts of N together with fixed organic carbon to  
215 northwestern bays. In line with this, we observed significant increases in TN from the lake  
216 center to Meiliang Bay and Gonghu Bay (Figure 2B), and  $\text{NO}_3^-\text{-N}$  and  $\text{NH}_4^+\text{-N}$  limitation in  
217 the overlying water during algal bloom period due to algae assimilation (Figure S2). The  
218 decomposition of this organic matter likely releases labile carbon substrates and consumes  
219 oxygen, creating favorable conditions for DNF in sediments. Low oxygen and availability of  
220 carbon result in converting algae concentrated zone into the hotspot of N loss by DNF. Due to  
221 frequent disturbances induced by winds, water column is difficult to be anoxia (Gao et al.,  
222 2017). We argue oxygen penetrated into sediments controls denitrification in sediments. Since  
223 the microelectrode needle is easily broken by hard stones in fields, we conducted incubation  
224 experiment to analyze responses of sediment oxygen profile to algal biomass, and found that  
225 the settlement of algal biomass can decrease oxygen penetration depth quickly in the surface  
226 sediment (Figure S3). In contrast, in the area where algae have left, sediment N loss was week

227 due to the limited availability of necessary substrate, namely  $\text{NO}_3^-$ -N and labile organic carbon.  
228 Thus, sediment N loss rates generally showed an increasing pattern from the lake center to the  
229 algae concentrated bays (Figure 3), and a similar temporal pattern with algal biomass in the  
230 lake center, where the rate was slightly lower during non-algal bloom period than the algal  
231 bloom period (Figure 4). Bernhardt (2013) outlined a paradigm where lakes need to be dirty  
232 (rich in nutrients) in order to support algal biomass at levels sufficient to fuel DNF for N  
233 removal. In shallow lakes, it is argued that wind-driven algal migration can effectively  
234 concentrate algal biomass into a 'dirty zone' where N removal might be maximized. Large  
235 amounts of N assimilated into the tissues of algal biomass were transported to this N removal  
236 region to be efficiently denitrified.

237 As a saying goes "A little wind kindles, much puts out the fire". If algae accumulate into dense  
238 mats, the decay of excessive algal biomass can deplete oxygen and limit oxygen penetration  
239 into sediments, causing severe anoxia (Karim et al., 2002; Zhu et al., 2020). Sediment  
240 incubation experiment indicated surface sediment suffered a longer anoxia period after more  
241 algal biomass settled (Figure S3). In the field, severe anoxia in bed sediments may persist for  
242 the entire algal bloom season due to persistent wind-driven algal biomass inputs. This persistent  
243 anoxia can inhibit NF and cause  $\text{NO}_3^-$ -N limitation on DNF, causing a break in coupled NF-  
244 DNF. We observed this to occur, and found that the effect was enhanced as algal biomass  
245 increased, suggesting that at high concentrations of algae,  $\text{NO}_3^-$ -N likely limited DNF. This  
246 explains higher sediment N loss rate during the non-algal bloom period than the algal bloom  
247 period in lake bays. During the algal bloom season, nitrifiers and denitrifier abundances were

248 relatively poor in the algae concentrated Meiliang Bay and Gonghu Bay (Figure 6). In the algae  
249 concentrated Meiliang Bay and Gonghu Bay, sediment N loss rates were negatively correlated  
250 with algal biomass (Figure 4D), and exhibited an opposite pattern with algal biomass as  
251 compared to the center of the lake (Figure 4B). The N loss rates in the lake bays were even  
252 lower in the hot summer than those in the cool winter, even lower than those in the algae-poor  
253 lake center (Figure 3). In sediments, direct DNF may also contribute to N loss, which is  
254 supposed to be enhanced by rich organic carbon and anoxia in the algae concentrated zone.  
255 Thus, we attributed the inhibited N loss to the breaking of the coupled NF-DNF. As anoxia  
256 receded post algal blooms, the coupled NF-DNF restarts, and the negative impacts of excessive  
257 algal biomass can weaken or disappear later, as seen in the increased N loss rate in the fall  
258 months (Figure 3). In contrast, in the lake center where algae have left, sediment N loss rates  
259 were positively correlated with algal biomass (Figure 4C) and showed a similar pattern with  
260 algal biomass (Figure 4A), suggesting that carbon availability may limit DNF under these  
261 conditions. Piecewise linear regression analysis indicated that, when algal biomass is over 70  
262  $\mu\text{g/L}$ , sediment N loss will be inhibited during algal bloom seasons in this study (Figure 5).  
263 Although the threshold may change in association with field conditions (temperature,  
264 turbulence intensity, etc.), it is possibly that, in portions of this eutrophic lake where wind-  
265 driven algal accumulations occur, water quality recovery is temporally prevented by the  
266 breaking of coupled NF-DNF, causing the redistribution of N removal amount over the whole  
267 year.

268 N biogeochemical cycles in lakes are complex, and many factors affect N loss in bed sediments,  
269 including zoobenthos and catchment features (Bonaglia et al., 2014; Nizzoli et al., 2018). In  
270 this study, we demonstrated that wind driven algal migration in the overlying water plays a key  
271 role in shaping sediment N loss patterns in shallow lakes. Algal migration created N loss  
272 hotspot in the algae concentrated zone. However, sediment N loss in this zone was inhibited  
273 during algal bloom seasons, and exhibited a negative relationship with algal biomass in the  
274 overlying water, which we associate with a broken link between NF and DNF. In the zone  
275 where algae have left, sediment N loss rates were relatively low and showed an increasing trend  
276 as algal biomass in the overlying water increased, suggesting an effective yet slow N removal  
277 process. By contrast, in deeper lakes, a larger fraction of algal biomass can be decomposed  
278 during the relatively long sinking time (Tammeorg et al., 2013). This may potentially result in  
279 algal blooms having a relatively smaller impact on sediment N processing in larger lakes,  
280 although it may also be that the associated impacts simply occur instead in the overlying water,  
281 a subject deserving of further study.

## 282 **5. Conclusions**

283 In this study, we investigated sediment DNF patterns along algal migration pathway in Taihu  
284 Lake, a shallow and eutrophic lake in China. The main findings are as follows:

- 285 (1) Wind driven algal migration in the overlying water manipulated DNF patterns in sediments  
286 in shallow lakes.
- 287 (2) A N loss hotspot in sediments was created in the algae concentrated zone, where N loss  
288 was, however, temporarily inhibited during algal bloom seasons and generally exhibited a

289 negative relationship with algal biomass.

290 (3) In the zone where algae have left, sediment N loss rate was relatively low and positively  
291 correlated with algal biomass.

292 (4) There is an algal biomass threshold. The sediment DNF rate increases beyond the threshold  
293 and decreases after the threshold.

294

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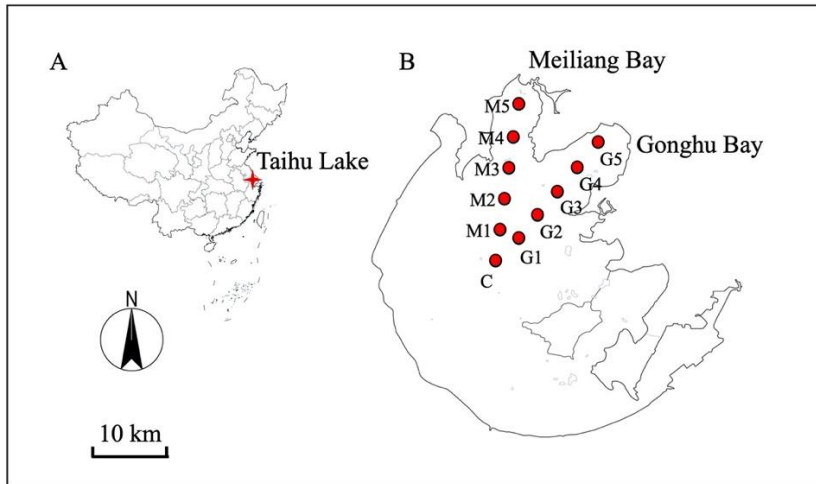
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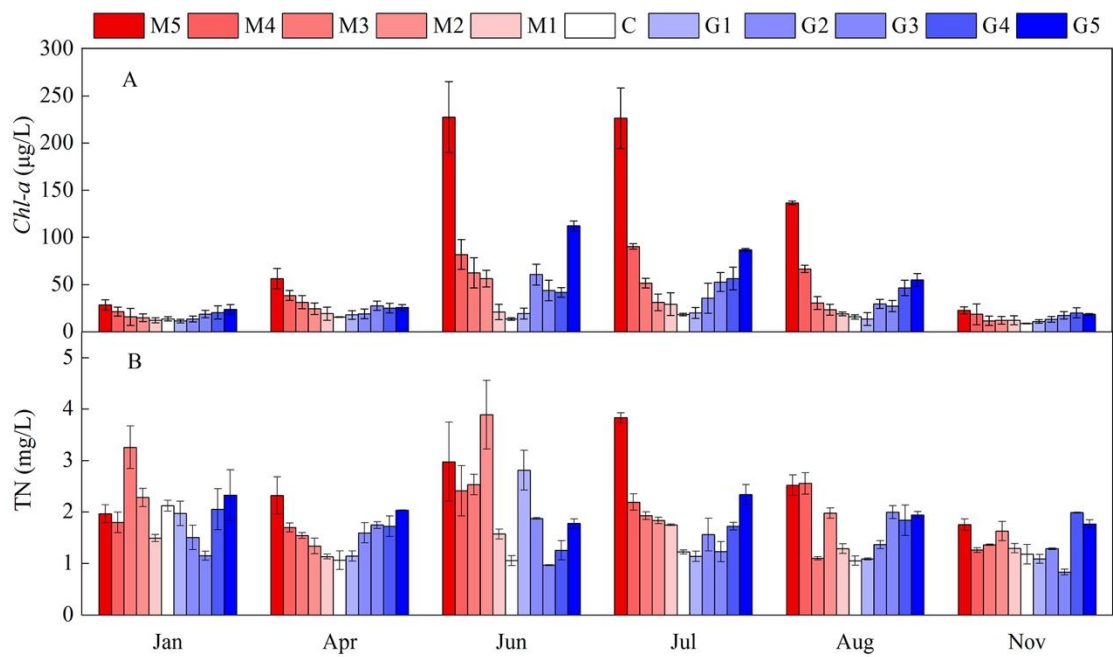


431

432 **Figure 1.** Map of Taihu Lake and locations of sampling sites in this study. Insert shows the

433 location of Taihu in China.

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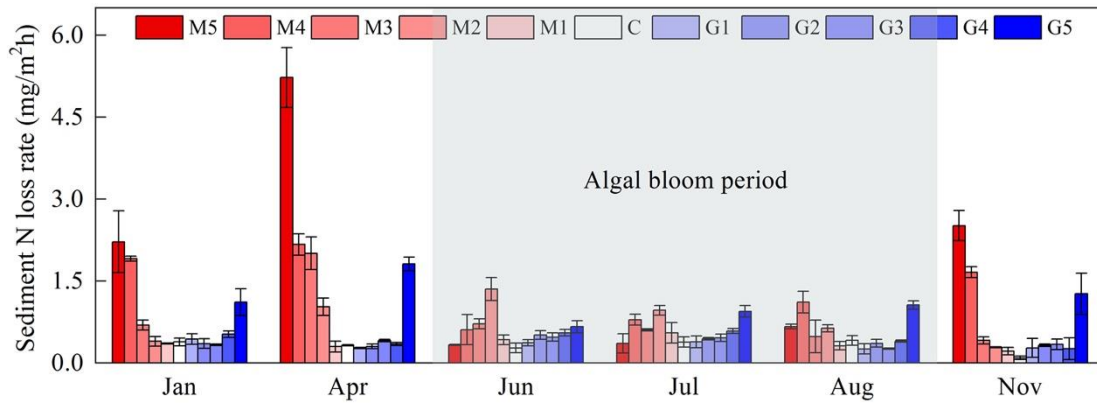
435

436 **Figure 2.** Temporal and spatial patterns of *Chl-a* and TN in the surface water along wind-

437 driven algal migration pathway. (A) *Chl-a*; (B) TN. Error bars indicate standard deviations.

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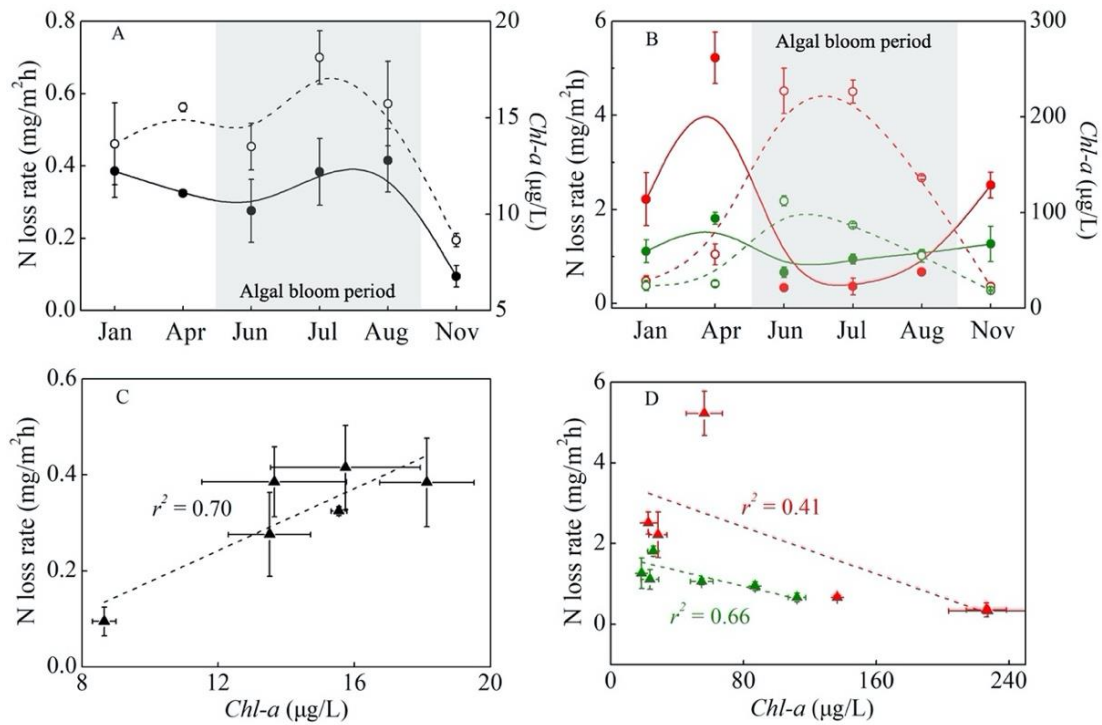


439

440 **Figure 3.** Temporal and spatial patterns of sediment N loss rate along algal migration pathway.

441 Error bars indicate standard deviations.

442



443

444 **Figure 4.** N loss pattern in bed sediments and its relationship with algal biomass in the

445 overlying water. In figure 4A and 4B, the black, red and olive circles indicate lake center,

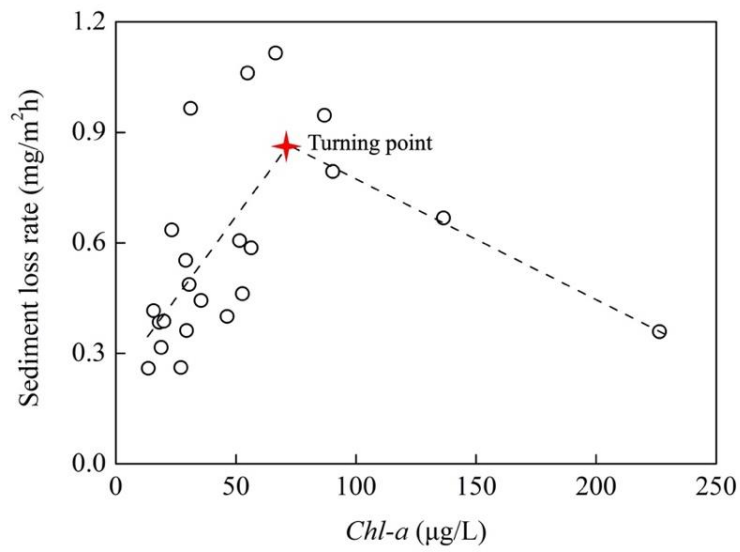
446 Meiliang Bay and Gonghu Bay, respectively, of which solid circles indicate N loss rate and

447 open circles indicate *Chl-a*. The lines were obtained using the B-spline model. In figure 4C and

448 4D, the black, red and olive triangles indicate lake center, Meiliang Bay and Gonghu Bay,

449 respectively. Error bars indicate standard deviations.

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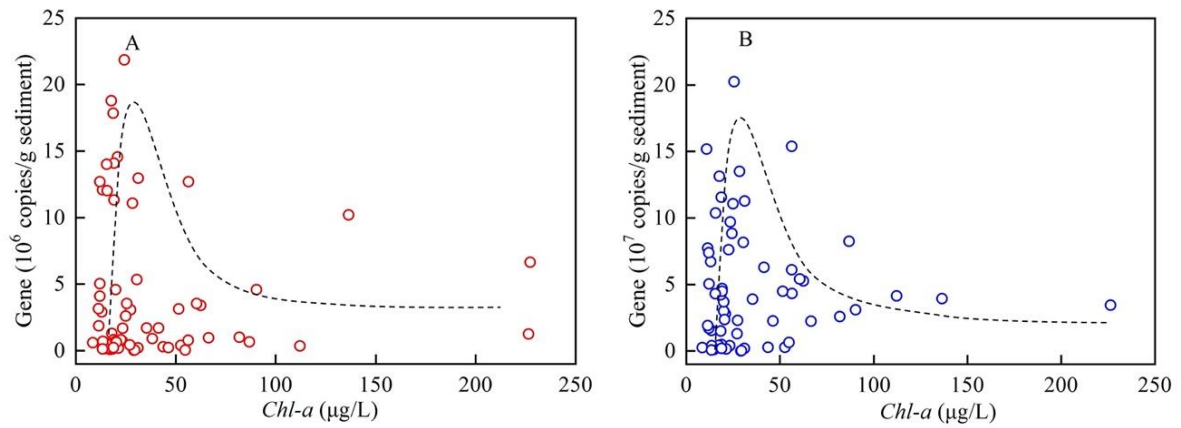
451

452 **Figure 5.** Changes of N loss rate in sediments as a function of algal biomass in the overlying

453 water. To avoid the disturbances caused by temperature, data were from all of sampling sites

454 in July and August, when water temperature showed no significant differences (Table S1).

455



456

457 **Figure 6.** The relationship between nitrifier and denitrifier abundances in sediments and *Chl-*  
 458 *a* in the overlying water. (A) Nitrifiers; (B) Denitrifiers. The abundance of nitrifiers was the  
 459 sum of *amoA* and *arch-amoA*, and the abundance of denitrifiers was the sum of *nirS*, *nirK* and  
 460 *nosZ* genes.

461